

# Organic matter formation from formaldehyde on Mars through time

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## Introduction:

Present-day Mars is extremely cold and dry, but geomorphological and geochemical evidence indicates a warm environment compatible with surface liquid water on early Mars. In addition to water, the existence of genomic and catalytic bio-molecules that support the fundamentals of life is critical for habitability. The Sample Analysis at Mars instrument onboard the Curiosity rover discovered organic matter in the early Martian sediment [1]. Moreover, a recent analysis discovered highly variable and <sup>13</sup>C-depleted carbon isotopic values in its organic [2]. However, its origin and the availability of biologically important molecules on early Mars remain uncertain.

Formaldehyde (H<sub>2</sub>CO), a highly soluble and reactive molecule, can be formed through photochemical reactions in planetary atmospheres. Dissolved H<sub>2</sub>CO in water would subsequently form diverse organic matter, including bio-important molecules, such as sugars and amino acids, by the formose-type reaction [3]. Moreover, H<sub>2</sub>CO can be generated from <sup>13</sup>C-depleted CO derived from CO<sub>2</sub> photolysis in the atmosphere [4]. The deposition of this photochemically produced H<sub>2</sub>CO may explain the origin of the strong isotopic depletion observed in Martian organic matter. Therefore, investigating the atmospheric synthesis of H<sub>2</sub>CO on early Mars is crucial for understanding prebiotic chemistry toward the potential of ancient Martian life. In this presentation, we summarize the production of H<sub>2</sub>CO and its carbon isotope composition on Mars through time. Finally, the origin of Martian organic matter will be discussed.

**Atmospheric production of formaldehyde through time:** First, we investigated the global average atmospheric production and deposition of H<sub>2</sub>CO in a 2-bar CO<sub>2</sub>-dominated atmosphere with H<sub>2</sub> and CO on early Mars at ~4–3 billion years ago (Ga). We adapted a 1D photochemistry model, PROTEUS (Photochemical and RaiatiOn Transport model for Extensive Use) [5], for early Martian conditions. We utilized the temperature and H<sub>2</sub>O vapor density profiles from the global mean results computed by a 3D Paleo-Mars Global Climate Model (PMGCM) [6] for a 2-bar CO<sub>2</sub>-dominated atmosphere with 0, 3, and 6 % H<sub>2</sub> and an obliquity of 40°. Our results show that the deposition flux of H<sub>2</sub>CO reaches the order of 10<sup>8</sup> or 10<sup>9</sup> cm<sup>-2</sup> s<sup>-1</sup> under conditions where the mixing ratio of H<sub>2</sub> is higher than 0.1%, regardless of the CO mixing ratio [7]. In a warm climate under a 6% H<sub>2</sub>

condition, the global average deposition flux of H<sub>2</sub>CO into the ocean is 3 × 10<sup>9</sup> cm<sup>-2</sup> s<sup>-1</sup>. Given the previously reported conversion rate from H<sub>2</sub>CO to ribose [3], the calculated H<sub>2</sub>CO deposition flux suggests a continuous supply of bio-important sugars on Noachian and early Hesperian Mars. After the warm climate periods, volcanic degassing would have decreased from the late Hesperian to the Amazonian, decreasing the mixing ratio of H<sub>2</sub> in the atmosphere [8]. Thus, the late Hesperian to early Amazonian was a transitional period from a high to a meager H<sub>2</sub>CO deposition flux (Figure 1).

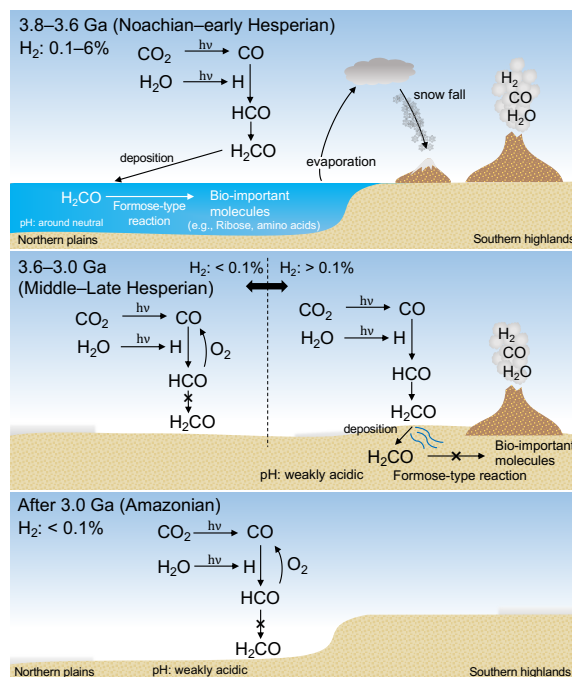


Figure 1. Scenario for the atmospheric H<sub>2</sub>CO production on Mars through time [7].

Second, we focused on a warm climate and calculated the global distribution of H<sub>2</sub>CO deposition across the globe in a PMGCM by creating a lookup table of H<sub>2</sub>CO production for key atmospheric parameters: surface temperature, surface pressure, surface H<sub>2</sub>O vapor, and UV flux. We assumed a 2-bar CO<sub>2</sub> atmosphere with 6% H<sub>2</sub> and 1% CO. As a result, the H<sub>2</sub>CO deposition flux is higher in regions with high concentrations of H<sub>2</sub>O vapor and efficient liquid water precipitation, corresponding to the northern ocean and its coastline at low latitudes.

### Carbon isotopic ratio in H<sub>2</sub>CO through time:

Third, we developed a coupled 1D photochemistry-climate evolution model incorporating carbon isotope fractionation processes induced by CO<sub>2</sub> photolysis, carbon escape to space, and volcanic outgassing in an early Martian atmosphere of 0.5–2 bar, composed mainly of CO<sub>2</sub>, CO, and H<sub>2</sub> to track the evolution of the carbon isotopic composition of C-bearing species [9]. The calculated carbon isotopic ratio in H<sub>2</sub>CO shows a significant depletion in <sup>13</sup>C with a minimum <sup>13</sup>C value of approximately −200‰, resulting from the isotopic fractionation induced by CO<sub>2</sub> photolysis (Figure 2). This ratio varies with changes in the background atmospheric conditions, such as CO/CO<sub>2</sub> ratio, surface pressure, and H<sub>2</sub> outgassing rate. Conversely, CO<sub>2</sub> becomes enriched in <sup>13</sup>C, consistent with an estimate from the carbonates preserved in the Martian meteorite Allan Hills 84001. These findings imply that certain amounts of the organic matter containing strongly depleted <sup>13</sup>C in the early Martian sediment measured by the Curiosity rover could have originated from the photochemically produced H<sub>2</sub>CO, undergoing subsequent condensation processes in water, such as formose-type reactions, during transient melting events in the late Noachian to Hesperian periods. Mixing with other organic sources, such as meteoritic material with higher δ<sup>13</sup>C values, could explain the observed variable δ<sup>13</sup>C values in Martian organic matter.

**Conclusion:** Our findings suggest that a weakly reducing early Martian atmosphere was suitable for H<sub>2</sub>CO photochemical production, which could subsequently convert into complex organic matter, including life’s building blocks, by formose reaction in water. Moreover, we showed that the carbon isotopic ratio of organic matter originating from photochemically produced H<sub>2</sub>CO are consistent with the measurement by the Curiosity rover. These results indicate that early Mars not only retained surface liquid water but also likely hosted organic compounds such as ribose, a prebiotic precursor originating from H<sub>2</sub>CO.

**References:** [1] Eigenbrode et al. (2018) *Science*, 360, 1096-1101 [2] House, C. H. et al. (2022) *Proc. Natl. Acad. Sci. U. S. A.* 119, e2115651119. [3] Ono et al. (2024) *Astrobiology*, 24, 489-497 [4] Ueno, Y. et al. (2024) *Nat. Geosci.* 17, 503–507. [5] Nakamura, Y. et al. (2023a) *Earth Planets Space* 75, 140. [6] Kamada et al. (2021) *Icarus*, 368, 114618. [7] Koyama, S. et al. (2024a) *Sci. Rep.* 14, 2397. [8] Grott et al. (2011) *Earth Planet. Sci. Lett.* 308, 391–400. [9] Koyama et al. (2024b) *Sci. Rep.*, 14, 21214.

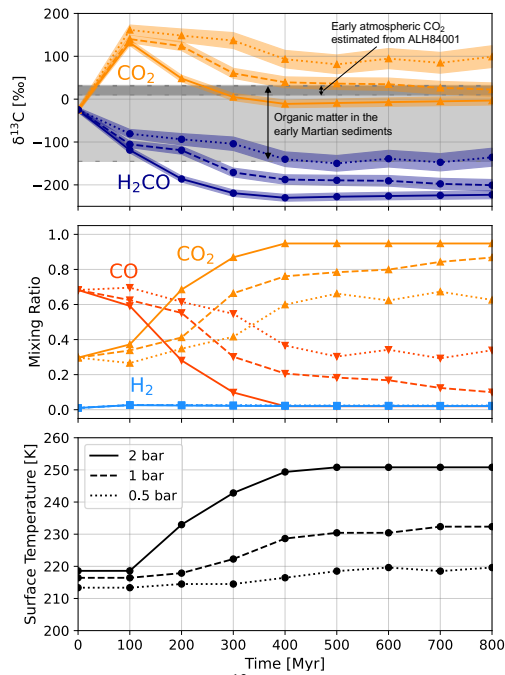


Figure 2. Evolution of δ<sup>13</sup>C (upper panel) in CO<sub>2</sub> and H<sub>2</sub>CO, mixing ratios of CO<sub>2</sub>, CO, and H<sub>2</sub> (middle panel), and surface temperature (lower panel) for 0.5-bar (dotted line), 1-bar (dashed line), and 2-bar (solid line) atmosphere, starting from 4 Ga [9].